

1. Motion of the Bay Water caused by Seismic Sea Waves.

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It is well known¹⁾ that when a bay, of which the depth is $h(x)$, the breadth $b(x)$, the length l , is subjected to the disturbance $\zeta = f(t)$ at its mouth $x = 0$, the elevation of the sea surface $\zeta(x, t)$ in the bay is given by

$$\zeta = gb(0)h(0) \sum_{s=1}^{\infty} \left(\frac{d\chi_s}{dx} \right)_0 \frac{\chi_s}{\lambda_s} \int_0^t f(\tau) \sin \lambda_s(t-\tau) d\tau, \quad (1)$$

where χ_s 's are normalized normal functions satisfying the equations:

$$\frac{d}{dx} \left[b(x)h(x) \frac{d\chi_s}{dx} \right] + \frac{\lambda_s^2}{g} b(x)\chi_s = 0, \quad (2)$$

$$\chi_s = 0 \quad \text{at} \quad x = 0, \quad (3)$$

$$d\chi_s/dx = 0 \quad \text{at} \quad x = l, \quad (4)$$

and

$$\int_0^l b(x)\chi_s^2 dx = 1. \quad (5)$$

In the case of a bay having a uniform depth H and breadth B , we have

$$\chi_s = \sqrt{2/Bl} \sin \left(s - \frac{1}{2} \right) \frac{\pi x}{l}, \quad \lambda_s = \frac{\sqrt{gH}}{l} \left(s - \frac{1}{2} \right) \pi, \quad (6)$$

so that

$$\zeta = \frac{2c}{l} \sum_{s=1}^{\infty} \sin \left(s - \frac{1}{2} \right) \frac{\pi x}{l} \int_0^t f(\tau) \sin c \left(s - \frac{1}{2} \right) \frac{\pi}{l} (t-\tau) d\tau, \quad (7)$$

where $c = \sqrt{gH}$. Writing that $t' = ct/l$, $\tau' = c\tau/l$, we have

1) Bull. Earthq. Res. Inst. Suppl. Vol. 1 (1934), 182.

$$\zeta = 2 \sum_{s=1}^{\infty} \sin\left(s - \frac{1}{2}\right) \frac{\pi x}{l} \int_0^{t'} f(l\tau'/c) \sin\left(s - \frac{1}{2}\right) \pi(t' - \tau') d\tau' \quad (7')$$

$$= \frac{2}{l} \sum_{s=1}^{\infty} \sin\left(s - \frac{1}{2}\right) \frac{\pi x}{l} \int_0^{t'} f(t - \xi/c) \sin\left(s - \frac{1}{2}\right) \frac{\pi \xi}{l} d\xi. \quad (7'')$$

Since we assume that $f(t)$ vanishes when $t < 0$, the upper limit of the integral in (7'') may be replaced by ∞ . Now divide the range of integration $0 \sim \infty$ into $0 \sim l$, $l \sim 2l$, $2l \sim 3l$, and change the variable ξ to ξ' , $2l - \xi'$, $2l + \xi'$, respectively. Then, noting that

$$f(x) = \frac{2}{l} \sum_{s=1}^{\infty} \sin\left(s - \frac{1}{2}\right) \frac{\pi x}{l} \int_0^{\infty} f(\xi) \sin\left(s - \frac{1}{2}\right) \frac{\pi \xi}{l} d\xi, \quad (8)$$

because $\sqrt{2/B} \sin\left(s - \frac{1}{2}\right) \pi x/l$ is a normal function, we obtain the following expression:

$$\begin{aligned} \zeta = & f(t-x/c) - f(t-2l/c-x/c) + f(t-4l/c-x/c) \\ & + f(t-2l/c+x/c) - f(t-4l/c+x/c) + \dots \end{aligned} \quad (9)$$

The same expression will be obtained more easily by the method of operational calculus. It can easily be seen from the above expression that ζ is composed of the direct wave and the waves reflected at the head and mouth of the bay. The coefficients of reflection are 1 and -1 in this case. When they are respectively α , ($-1 < \alpha < 0$) at the mouth and β , ($1 > \beta > 0$) at the head of the bay owing to such causes as the scattering of the energy, the friction at the bottom, etc., we have the following expression instead of (9):

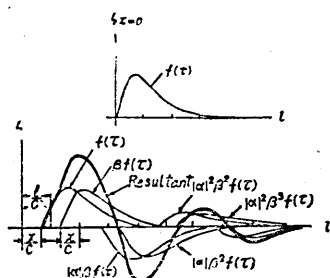


Fig. 1. $\zeta(t)$ at the mouth of the bay (above) and $\zeta(t)$ at any place in the bay (below), which is composed of the direct wave and reflected waves.

$$\begin{aligned} \zeta = & \sum_{n=0}^{\infty} \alpha^n \beta^n [f(t-2nl/c-x/c) \\ & + \beta f(t-2n+1)l/c+x/c)] \end{aligned} \quad (10)$$

It can be seen from (10) that in the bay there develop oscillations of the period $4l/c$ and their harmonics of odd orders, irrespective of the shape of $f(t)$.

It will be worthy of notice that from (10) we can deduce the following important relations :

$$\Phi(t) \equiv \zeta(t) - \alpha\beta\zeta(t-2l/c) \equiv f(t-x/c) + \beta f(t-2l/c+x/c), \quad (11)$$

and

$$f(t) = \sum_{n=0}^{\infty} (-1)^n \beta^n \Phi(t-2nl-x/c+x/c). \quad (12)$$

By means of this expression, we can calculate the original shape of the seismic sea waves from the observation made at any point in the bay, provided that α and β are known. In most cases β is nearly equal to unity. As for α , it depends on the shape of the bay, but we can estimate it from the period-amplitude diagram of the record of the seismic sea waves. Assuming, for the sake of simplicity, that $\beta = 1$, $x = l$, and $f = e^{ipt}$, we have

$$|\zeta| = 2/\sqrt{(1-\alpha)^2 \cos^2 pl/c + (1+\alpha)^2 \sin^2 pl/c}, \quad (13)$$

so that the ratio $|\zeta|_{\max}/|\zeta|_{\min}$ is equal to $(1-\alpha)/(1+\alpha)$.

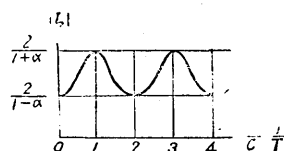


Fig. 2. Theoretical period-amplitude diagram, from which α may be estimated.