博士論文 (要約)

Studies on the estimation of water temperature experienced during the larval stage of Pacific bluefin tuna *Thunnus orientalis* using SIMS oxygen isotope analysis of otoliths

(SIMSによる耳石酸素安定同位体比分析を用いた クロマグロ仔魚期の経験水温の推定に関する研究)

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Chapter 1 cannot be viewed because the contents will be published in an academic journal.

Chapter 2 cannot be viewed because the contents will be published in an academic journal.

Temperature reconstruction using otolith $\delta^{18}O$

3.1. Introduction

Ocean warming causes significant impacts on marine species and ecosystems, including high mortality, distribution shifts, and loss of spawning and nursery habitats (Perry et al., 2005, Kimura et al., 2010, Muhling et al., 2011, 2015). Species that spawn seasonally in relatively limited areas are particularly vulnerable to increasing water temperature, as their optimum range in spawning temperatures tends to be restricted. Pacific bluefin tuna is a highly migratory species that spawns in waters near the Nansei Islands in the western North Pacific from May to June and in the Sea of Japan from July to August (Yonemori 1989, Ohshimo et al., 2017). Adult fish in the western North Pacific spawn at temperatures between 26 and 29°C, whereas those in the Sea of Japan initiate spawning at temperatures greater than 20°C (Chen et al., 2006, Tanaka 2011, Suzuki et al., 2014, Okochi et al., 2016). In the laboratory, the growth rate and survival of Pacific bluefin tuna larvae significantly decrease when temperatures exceed 29°C (Kimura et al., 2007). In fact, projected temperature in the current spawning sites is expected to increase by more than 3°C by 2100 under the most extreme IPCC climate-warming scenario (IPCC 2007) and become unsuitable for Pacific bluefin tuna to spawn (Kimura et al., 2010). As bluefin larvae are particularly vulnerable to thermal stress, warming sea temperatures are likely to cause significant impacts on their early growth and survival. However, the effects of ongoing climate change on the early life stages of Pacific bluefin tuna are poorly understood due to a lack of empirical evidence and methods to study such effects.

Oxygen isotope ratios (δ^{18} O) in otoliths, biogenic calcium carbonate (aragonite) found in the inner ear of teleost fish (ray-finned bony fish), has been widely used as a natural tag to reconstruct water temperatures and salinity conditions experienced by fish (Thorrold et al., 1997, Campana 1999, Jones & Campana 2009). Such reconstructions of past environment are possible because otoliths generally develop at or close to the isotope equilibrium with ambient water, and many studies have demonstrated the temperature-dependency of otolith δ^{18} O for various fish species under laboratory conditions (Kalish 1991, Thorrold et al., 1997, Høie et al., 2004, Kitagawa et al., 2013). Existing methods of temperature reconstruction for fish mostly rely on $\delta^{18}O_{\text{otolith}}$ measurements by conventional isotope ratio mass spectrometer (IRMS), which often involves a milling process to obtain a relatively large amount of otolith powder for analysis (usually > a few tens of micrograms with a minimum weight requirement of 15 µg). Ambient water temperatures have previously been reconstructed using IRMS for sockeye salmon *Oncorhynchus nerka* (Zazzo et al., 2006), alewife *Alosa pseudoharengus* (Dufour et al., 2008), Atlantic cod *Gadus morhua* (Jones & Campana 2009, Leesen et al., 2020), turbot *Scophthalmus maximus* (Imsland et al., 2014), and chub mackerel *Scomber japonicus* (Higuchi et al., 2019), most of which have a monthly to annual resolution depending on the otolith size. The limited temporal resolution due to the sample mass requirement of the IRMS is inevitable and makes it particularly difficult when analyzing the otolith core and edge.

The recent developments in secondary ion mass spectrometry (SIMS) δ^{18} O analysis of otoliths have enabled a high-resolution reconstruction of migration and life history characteristics of marine species (Hanson et al., 2010, Matta et al., 2013, Shiao et al., 2014, Helser et al., 2018a, Shirai et al., 2018, Willmes et al., 2019). Unlike conventional IRMS, SIMS is capable of determining isotopic composition within a spatial resolution of 5–15 µm, which allows sub-annual, seasonal, and even weekly or much shorter timescale analyses with high accuracy and precision (Valley & Kita 2009, Kita et al., 2009). While a recent study (Sakamoto et al., 2019) reconstructed migration histories of an individual Japanese sardine *Sardinops melanostictus* with 10–30 days resolution (20–30 days around the core regions and 10–15 days toward the edge) using microvolume isotope analysis measured by continuous-flow IRMS (CF-IRMS), SIMS provides even finer temporal resolution, particularly for the otolith core and edge. High-resolution reconstruction of experienced temperatures using SIMS $\delta^{18}O_{otolith}$ has important utility to investigate the early life history of fish that may be affected by increasing water temperature associated with climate change.

In this chapter, we aim to develop a method to reconstruct ambient water temperatures experienced during the larval period of an individual fish using SIMS $\delta^{18}O_{\text{otolith}}$ analysis. $\delta^{18}O_{\text{otolith}}$ of five adult Pacific bluefin tuna are measured from the otolith core to edge by SIMS, and the measured SIMS $\delta^{18}O_{\text{otolith}}$ values were compared to those measured by CF-IRMS. Water temperatures are then estimated using a temperature-dependent oxygen isotope fractionation equation for Pacific bluefin tuna larvae that has been already established in the previous study (Kitagawa et al., 2013). The temperature reconstruction technique presented here allows for a high-resolution investigation of the early life history of fish and provides a more thorough understanding of the characteristics of survivors and thermal environment that may constrain their early growth and survival.

3.2. Materials and Methods

3.2.1. Sample collection and sample preparation

Otoliths were collected from the defrosted heads of Pacific bluefin tuna caught mostly by longline fishing gear in waters around Japan from 2017 to 2018 (Fig. 3-1). In total, otoliths from 119 bluefin tuna (sub-adult and adult) were collected. As data on fish weight was only available, the fork length of fish was estimated using a weight–length relationship established by Kai (2007). The otolith weight (length) and estimated weight (fork length) of the fish collected in this study and those of young-of-the-year fish (Suzuki, 2019) were measured, and their relationship with fish weight and fork length was determined, respectively (Figs. 3-2, 3-3, 3-4).

Age of the fish was determined according to the age reading protocol that has been already developed for Pacific bluefin tuna (Shimose & Ishihara, 2015) (Fig. 3-5). The relationship between

otolith weight (length) and fish weight for sub-adult and adult, young-of-the-year fish, and those combined is shown in Fig. 3-2a (b), Fig. 3-3a (b), and Fig. 3-4a (b), respectively.

Five adult fish samples were randomly selected and one of paired otoliths from each individual was used for SIMS δ^{18} O analysis. Catch location, date of catch, and biological information of bluefin tuna samples are shown in Fig. 3-6 and Table 3. Otoliths were cleaned and rinsed with double deionized water (Milli-Q water) to remove any remaining muscle tissues, air-dried in a clean environment, and stored in microtubes for later analyses. For CF-IRMS δ^{18} O analysis, the three otolith sections that were analyzed by SIMS were used.

For the preparation of otolith samples for SIMS δ^{18} O analysis, we used the sample preparation protocol for a single otolith thin section that has been previously developed in Chapter 2. A detailed sample preparation protocol can be found in Fig. 2-1. Briefly, an otolith was mounted on a microscope slide with thermoplastic cement perpendicular to the sagittal plane with its sulcus side facing down. The otolith core was observed under an inverted microscope (IX-71, Olympus) and straight lines were drawn on the glass slide at 250 to 270 µm on each side of the otolith core using waterproof ink and a comic pen. An otolith thin section was cut out together with the glass slide along the drawn lines with an automatic low-speed precision cutter (IsoMet 5000, Buehler) equipped with a 0.3 mm thick diamond blade (IsoMet 15LC, Buehler). The sectioned otolith was then removed from a strip of glass by rinsing it with acetone and was allowed to air-dry in a laminar flow hood. The sectioned otolith was fixed in the center of a 2.54 cm silicon mold and embedded in epoxy (EpoxyCure 2 Resin, Buehler) along with a UWC-3 standard, that was placed right above the otolith. It was then kept at room temperature for 24 h to cure the resin. The epoxy disk containing an otolith thin section and a piece of standard material was ground with a grinding machine equipped with 70 and 13 µm diamond cup wheels (Discoplan-TS, Struers) until the distance from the otolith surface to the core reached 15 to 20 μ m. It was then successively polished using 6, 3, and 1 μ m diamond pastes on a fine grinding disc (MD-Largo, Struers) to expose the core on a flat mirror-finished surface.

Before analysis, the samples were cleaned in an ultrasonic cleaner and dried in a vacuum oven at 40 °C for 2 h. They were then sputter-coated with ~60 nm gold.

3.2.2. SIMS δ^{18} O analysis

Otolith oxygen isotope ratios were measured *in situ* using a CAMECA IMS 1280-HR large radius, multi-collector ion microprobe (SIMS) at the Kochi Institute for Core Sample Research, Japan Agency for Marine-Earth Science and Technology (JAMSTEC). Five otolith thin sections were prepared for SIMS $\delta^{18}O_{\text{otolith}}$ analyses using the sample preparation protocol developed in the present study (Fig. 2-1). The $\delta^{18}O_{\text{otolith}}$ values were measured from otolith core to edge along the growth axis for each otolith sample (Fig. 3-7).

The SIMS analytical conditions that were used for $\delta^{18}O_{\text{otolith}}$ measurements in this study are described in detail in Kita et al., (2009). The sample surface was sputtered by a 20 kV accelerated ¹³³Cs⁺ primary ion beam of 1.5–1.8 nA focused to a diameter of 10 to 15 µm, resulting in a pit of ~1 µm depth (Fig. 3-7b). The secondary ions (¹⁶O⁻, ¹⁸O⁻, and ¹⁶OH⁻) were accelerated at 10 kV and detected simultaneously by three Faraday cup detectors. Since hydrogen is present in the SIMS chamber even under ultra-high vacuum conditions, measured ¹⁶OH^{-/16}O⁻ ratios were background-corrected by subtracting the average ¹⁶OH^{-/16}O⁻ ratio of the UWC-3 standard (nominally anhydrous minerals) bracketing analyses from the ¹⁶OH^{-/16}O⁻ ratio of $\delta^{18}O_{\text{otolith}}$ measurements. The background-corrected ¹⁶OH^{-/16}O⁻ ratios served as a proxy for the relative hydrogen content contained in otolith samples. Each analysis took ~3 min, consisting of pre-sputtering (10 s), automatic centering of the secondary ion beam (90 s), and the isotopic measurements with 20 analytical cycles (40 s). The count rates for ¹⁶O⁻ and ¹⁸O⁻ were 1.7–2.5 × 10⁹ and 3.5–5.1 × 10⁶ counts per second (cps), respectively.

For accurate calibration of SIMS δ^{18} O measurements in biogenic carbonate samples, a homogeneous biocarbonate standard with a matched-matrix is needed. The UWC-3 standard is a

chemically and isotopically homogeneous calcite standard which has a similar chemical composition as otoliths (aragonite), and thus all $\delta^{18}O_{\text{otolith}}$ measurements were normalized with this standard in our study (Kozdon et al., 2009). Every 10 to 15 unknown sample measurements were bracketed by 10 analyses of a UWC-3 calcite standard (5 analyses before and after each group of unknown samples) to calculate the spot-to-spot precision of sample analyses and to correct for instrumental mass fractionation. The precision of sample analyses for all 5 otolith thin sections was ±0.3 to ±0.6‰ (2 standard deviations).

After analysis, each spot was observed on scanning electron microscope (SEM) images taken with an electron probe micro analyzer (JXA-8230, JEOL) to check for any cracks and inclusions that might bias the resulting δ^{18} O values (e.g. Weidel et al., 2007). No spots had such surface irregularities. In addition, the secondary ion yield (¹⁶O⁻, cps/nA) relative to the mean of the UWC-3 standard bracketing analyses were used to assess the quality of each spot measurement and check for any extreme outliers. Raw SIMS δ^{18} O measurements are presented in Table S1.

For comparison purposes, all $\delta^{18}O_{\text{otolith}}$ values were converted from VSMOW to VPDB by using the latest published conversion equation ($\delta^{18}O_{\text{VSMOW}} = 1.03092 * \delta^{18}O_{\text{VPDB}} + 30.92$, Brand et al., 2014, Kim et al., 2015).

3.2.3. Microvolume δ^{18} O analysis by CF-IRMS

Microvolume CF-IRMS δ^{18} O analysis was conducted to compare δ^{18} O_{otolith} values measured by SIMS and CF-IRMS. Three otolith thin sections that were analyzed by SIMS were used for microvolume CF-IRMS δ^{18} O analysis. δ^{18} O analyses were performed with an IsoPrime100 isotope ratio mass spectrometer (Isoprime Ltd, Cheadle Hulme, UK) equipped with a customized continuousflow gas preparation system (MICAL3c) at National Institute of Technology, Ibaraki College, Japan. This system can measure isotope ratios of calcium carbonate samples with a minimum sample mass of 0.2 μ g (about 1/100 of the sample mass required for commercially available IRMS systems) with high precision and accuracy (Ishimura et al., 2004, 2008). By micromilling the otolith material deposited during the same growth period that is being analyzed by SIMS, it is possible to compare average SIMS and CF-IRMS $\delta^{18}O_{\text{otolith}}$ values.

A high precision micromilling system (Geomill326) was used for milling the specific regions of otolith samples. This system comprised a carbide bur fixed over an XYZ sample stage, a high-resolution camera, and a computerized image analyzer. An otolith image with marks indicating the target milling areas was imported into the system and milling paths were configured on the computer. The target milling areas were set in the otolith region where the measured SIMS $\delta^{18}O_{\text{otolith}}$ values were stable, and they covered roughly 3–4 SIMS beam spots (Fig. 3-8). First, an unwanted area right next to the target milling path was milled and removed from otolith to avoid cross-contamination. The removed otolith powders were also collected and used for analysis as supplementary samples to increase the dataset. For each otolith sample, two to three target paths were milled along the growth rings to obtain powder samples. The milled paths resulted 15 to 80 µm wide, 250 to 350 µm long, and 60 to 90 µm deep.

The life stages of the corresponding milled paths were estimated based on the distance from otolith core and the location of the annual growth increments. Each path corresponded to either the juvenile (a few months old), immature (about 5 months old to age 1+), or sub-adult stage (age 2–3). The non-parametric Kruskal-Wallis test was used to determine the statistical difference in resulting offset values between SIMS and CF-IRMS among different life stages.

The amount of powder produced from each path was 0.9 to $3.5 \ \mu g$. The milled powder was carefully collected and placed on to a small piece of glass using a needle under a microscope, and then put into the bottom of a reaction tube. The aragonite powder was then reacted with phosphoric acid at 25°C, and the evolved CO₂ was purified in a stainless vacuum line. After further purification

using a helium-purged purification line, the purified CO₂ was introduced into the mass spectrometer. Samples that weighed more than 2.0 µg were analyzed twice. The results are reported in standard δ notation (‰) relative to Vienna Pee Dee Belemnite (VPDB). We measured a lab standard CO₂ gas for the determination of the analytical precision of pure CO₂ gas ($\delta^{13}C_{VPDB} = -1.56\%$ and $\delta^{18}O_{VPDB} = -4.42\%$, Nishida & Ishimura 2017) three or more times every day. The analytical precision was better than ± 0.1‰ (± 1SD) for the entire analysis.

To compare the difference between SIMS and CF-IRMS values, the average SIMS $\delta^{18}O_{otolith}$ values were calculated by averaging the SIMS spot measurements adjacent to or within each milling path and then compared to CF-IRMS $\delta^{18}O_{otolith}$ value. A linear regression analysis was performed to determine the correlation between values measured by the two methods. In addition, the Wilcoxon signed rank test was applied to test the statistical difference between the average SIMS and CF-IRMS $\delta^{18}O_{otolith}$ values, and the Tukey's test was used to detect outliers. All statistical analyses were performed using the R software (version 3.3.2).

3.2.4. Temperature reconstruction for larval stages of adult bluefin tuna

Ambient water temperatures experienced during the larval stages of adult bluefin were estimated using SIMS $\delta^{18}O_{\text{otolith}}$ of otolith core region. The daily growth increments were counted around the core region on the SEM images. The SIMS spot measurements that were made within 20 days post hatch (DPH) were used for the temperature estimation as the larval period of bluefin tuna lasts for approximately 20 days. For an accurate estimation of ambient water temperatures, a species-specific fractionation equation and $\delta^{18}O_{\text{seawater}}$ are needed. To reconstruct ambient temperatures, we used the oxygen isotope fractionation equation for Pacific bluefin tuna larvae proposed by Kitagawa et al. (2013):

$$\delta^{18}O_{\text{otolith (VPDB)}} - \delta^{18}O_{\text{water (VPDB)}} = 5.193 - 0.270 \times T$$
 (3-1)

where $\delta^{18}O_{water}$ is the $\delta^{18}O$ value of ambient water, and *T* is the water temperature in °C. For $\delta^{18}O_{water}$ in the equation, we applied the $\delta^{18}O$ value of +0.22‰ (VSMOW) which is the average $\delta^{18}O_{water}$ of the main spawning ground around the Nansei Islands (123.56-131.00°E, 24.04-26.09°N) in May to June from 2008 to 2010 (Uozato 2011). The $\delta^{18}O_{seawater}$ value was corrected on the VPDB scale by simply subtracting 0.27‰ (Hut 1987). The mean temperatures were calculated for the samples that had multiple $\delta^{18}O_{otolith}$ measurements made around the core regions. Although the focus was on the reconstruction of ambient water temperatures experienced during the larval period of the fish, the lifetime temperature history was estimated to evaluate how well $\delta^{18}O_{otolith}$ records ambient water temperatures in the older stages of Pacific bluefin tuna.

3.3. Results

3.3.1. Comparison of SIMS and CF-IRMS $\delta^{18}O_{otolith}$ measurements

In total, 22 paths were milled by a micromilling system and $\delta^{18}O_{otolith}$ of collected powders from each milled path was measured by CF-IRMS. After carefully examining the accuracy of how well each milled path captured the same growth zone as the SIMS spots, 18 samples, including 9 main samples and 9 supplementary samples, were selected to assess the difference between SIMS and CF-IRMS $\delta^{18}O_{otolith}$ values. The average values of multiple CF-IRMS measurements were used when the milling accuracy of a single milling path was low (See Appendix 2 for a complete list of data). The precision of $\delta^{18}O_{otolith}$ values measured by CF-IRMS was better than that of SIMS.

The $\delta^{18}O_{\text{otolith}}$ values measured by CF-IRMS were significantly higher than those measured by SIMS (Fig. 3-9, Wilcoxon signed-rank test, p < 0.001) except for two measurements in which the CF-IRMS value was 0.12‰ and 0.47‰ lower than the average SIMS value, respectively. We considered

the CF-IRMS $\delta^{18}O_{\text{otolith}}$ measurement that resulted 0.47‰ lower than the average SIMS $\delta^{18}O_{\text{otolith}}$ value as an outlier based on the Tukey's outlier detection method. This measurement was taken from a relatively large area of the otolith deposited during the second to third years of life, where the interannual variation in water temperature is expected. It is likely that this temperature variation was not well reflected in the average SIMS value since only 3 spot measurements were averaged over more than a half-year period. Therefore, this measurement was excluded for calculating a linear regression curve and the average SIMS–CF-IRMS difference.

There was a significant positive correlation between SIMS and CF-IRMS $\delta^{18}O_{\text{otolith}}$ values ($r^2 = 0.79, p < 0.001$) with a slope of 1.1408 and an intercept of 0.0704. The slope is 1 within the 95% confidence interval (lower limit = 0.82, upper limit = 1.46) with a high correlation coefficient ($r^2 = 0.78$), and thus a linear regression with a slope of 1 was fitted to the data to calculate the y-intercept, which is the average difference (Fig. 3-10). The average offset between SIMS and CF-IRMS $\delta^{18}O_{\text{otolith}}$ values was 0.41‰, with SIMS yielding lower values. The SIMS–CF-IRMS $\delta^{18}O_{\text{otolith}}$ correction equation for Pacific bluefin tuna can be expressed as:

SIMS
$$\delta^{18}O_{\text{otolith (VPDB)}} = \text{CF-IRMS } \delta^{18}O_{\text{otolith (VPDB)}} - 0.41$$
 (3-2)

No significant difference was observed between $\delta^{18}O_{\text{otolith}}$ measurements at different life stages (Kruskal-Wallis test, p = 0.29), and thus the consistent application of this offset correction equation to all SIMS measurements was considered appropriate.

3.3.2. Seasonal variations in SIMS $\delta^{18}O_{otolith}$ profiles

In total, 259 $\delta^{18}O_{otolith}$ measurements were made on the otoliths of five Pacific bluefin tuna by SIMS (Table 3). The total number of spots measured per otolith was 42 to 78 (two life-history transect

lines were analyzed for T64R). The length of each transect ranged between 2.5 mm and 2.8 mm with a spot-to-spot distance ranging from 16 μ m to 153 μ m around the core region, and 26 μ m to 337 μ m toward the edge. The spatial resolution around the core region in one of the samples was more than 10 times higher compared to the conventional IRMS method previously used for Pacific bluefin tuna otoliths (Shiao et al., 2010). The temporal resolution of SIMS spots was 3–5 days near the core region and roughly a few weeks to a month on the outer edge depending on the age of fish.

The offset-corrected SIMS $\delta^{18}O_{\text{otolith}}$ values for 5 Pacific bluefin tuna samples are plotted in Fig. 3-11 (see left y-axis). High-resolution $\delta^{18}O_{\text{otolith}}$ profiles of all otolith samples showed distinct seasonal variations with an increasing trend from the otolith core to about 1250 µm. The average $\delta^{18}O_{\text{otolith}}$ values from the core to about 750 µm ranged between -3.1 to -2.5% (VPDB), and sharply increased toward the first annual increment (opaque zone), peaking at -1.3 to -0.4%. After the increase, the $\delta^{18}O_{\text{otolith}}$ values decreased, showing a cyclical pattern toward the edge fluctuating mostly between -2.5 to -1.5% (VPDB). The $\delta^{18}O_{\text{otolith}}$ of the core regions corresponding to the larval stages (10 to 20 DPH) of bluefin tuna ranged between -3.1 to -1.9% (VPDB).

The background-corrected ¹⁶OH^{-/16}O⁻ ratios measured for the five otoliths ranged between 0.017 and 0.031 and had a general inverse relationship with the SIMS $\delta^{18}O_{\text{otolith}}$ values (there is no impact of individual differences on this relationship) (Fig. 3-12a). This indicates that the relative hydrogen content in the otolith increases with lower SIMS $\delta^{18}O_{\text{otolith}}$ values. Overall, higher ¹⁶OH^{-/16}O⁻ ratios resulted in larger SIMS–CF-IRMS $\delta^{18}O_{\text{otolith}}$ differences (Fig. 3-12b), which is consistent with the inverse trend seen in Fig. 3-12a.

3.3.3. Estimation of temperature experienced during larval period

Core-to-edge water temperature profiles of all samples are shown in Fig. 3-11 (see right y-axis). The estimated temperatures experienced during the larval stages ranged between 26.7°C to 30.7°C among the individuals (T64R: $30.7 \pm 1.3^{\circ}$ C [~20 DPH], T75R: $27.9 \pm 1.0^{\circ}$ C [~10 DPH], T104L: 26.7 $\pm 1.0^{\circ}$ C [~12 DPH], T118R: $28.9 \pm 0.9^{\circ}$ C [~12 DPH], T131R: $28.4 \pm 1.4^{\circ}$ C [~12 DPH]). After the year 0, temperature raged mostly between 24 and 30°C, and never reached 35°C.

3.4. Discussion

An increasing number of experimental and modeling studies have shown significant impacts of projected climate change on the early growth and survival of various fish species (Kimura et al 2007, Pankhurst & Munday 2011, Moyano et al., 2017). Generally, fish larvae are more sensitive to temperature variations than juveniles and adults as they have narrower thermal tolerance ranges (Pörtner & Peck 2010, Moyano et al., 2017), making them particularly vulnerable to climate change. However, the effects of ongoing climate change-driven ocean warming on the early life history of fish remain largely unexplored for many species, mostly due to difficulties in monitoring long-term responses to climatic stressors. Thus, the advancement of techniques that can quantitatively estimate past environments actually experienced by fish during key phases of the life cycle is essential to fill this knowledge gap. In this study, a high-resolution temperature reconstruction technique using SIMS $\delta^{18}O_{otolith}$ analysis was developed and applied to Pacific bluefin tuna, a species of great economic importance, whose early larval growth and survival may be constrained by climate change.

The 0.41‰ offset observed between SIMS and CF-IRMS $\delta^{18}O_{\text{otolith}}$ values for bluefin tuna otoliths is most likely due to incomplete correction for "matrix effects" by SIMS methods. Matrix effects refer to an instrumental mass fractionation caused by different chemical compositions and structures between given samples and standard materials, which shifts measured values (Eiler et al., 1997, Riciputi et al., 1998, Śliwiński et al., 2016, 2017, Wycech 2018). Otoliths contain a small amount of organic proteins, namely otolith matrix protein-1 and Otolin-1 (Murayama et al., 2000, 2002). The presence of these proteins in the otolith would be responsible for a subtle change in instrumental mass fractionation, which results in lower SIMS $\delta^{18}O_{otolith}$ values relative to those measured by CF-IRMS and higher ¹⁶OH⁻ ion yields compared to that of the calcite standard. The general inverse relationship between ¹⁶OH⁻/¹⁶O⁻ ratios and SIMS $\delta^{18}O_{otolith}$ values (Fig. 3-12a) may be the result of the incorporation of more proteins into the otolith matrix which is thought to relate to fast growth in summer (or less proteins in winter due to slow growth). Water content (OH⁻) of the otoliths, if any, is also responsible for the lower $\delta^{18}O_{otolith}$ values. These organic proteins and water content bias SIMS $\delta^{18}O_{otolith}$ values because they are measured together with calcium carbonates, whereas they do not affect CF-IRMS $\delta^{18}O_{otolith}$ values since these proteins do not react with phosphoric acid at 25°C that is used in a digestion process to generate CO₂ gas. Furthermore, the systematic difference in isotopic fractionations caused by sputtering different crystalline structures (the biogenic aragonite samples and calcite standard) may contribute to the observed offset (Linzmeier et al., 2016). Although matrix effects are likely the primary cause of the SIMS–CF-IRMS difference, other potential factors (e.g. milling and roasting effects) may influence the measurement results of SIMS and CF-IRMS. The effects of roasting and other factors are discussed in detail in Wycech et al. (2018) which investigated the $\delta^{18}O$ difference between SIMS and IRMS using foraminiferal shells.

The offset of 0.41‰ found in this study is within the range of offset values previously reported for otolith of other fish species and biocarbonate samples. Orland et al. (2015) has reported that there is a consistent sample-dependent offset in δ^{18} O typically less than 1.0‰, but it can go up to 1.8‰ depending on different sample matrices including biocarbonates and speleothems. Matta et al. (2013) used roasting to remove organic materials and observed an offset of about 1‰ in SIMS δ^{18} O values between roasted and unroasted otoliths of a yellowfin sole *Limanda aspera*. Helser et al. (2018b) also observed a 0.5‰ offset between SIMS and CF-IRMS δ^{18} O_{otolith} measurements in otoliths of Pacific cod *Gadus macrocephalus*. The temperature estimation without the correction of the 0.41‰ offset using SIMS δ^{18} O_{otolith} values of Pacific bluefin tuna would cause a bias of 1.5°C in the estimates, resulting in some unrealistically high temperature estimates (i.e. >34°C). While it is unknown to what

extent organic proteins and water content contribute to an overall offset, it is necessary to determine these offsets between SIMS versus IRMS when estimating ambient water temperatures from SIMS $\delta^{18}O_{otolith}$ values for different species.

The SIMS analyses performed on five otolith samples of Pacific bluefin tuna revealed fine-scale $\delta^{18}O_{\text{otolith}}$ profiles from the core to the edge with clear seasonal variations (Fig. 3-11). As water temperature and $\delta^{18}O_{\text{otolith}}$ are negatively correlated (Devereux 1967, Høie et al., 2004, Thorrold et al., 1997), the initial increase of the $\delta^{18}O_{\text{otolith}}$ values observed toward the first annual increment (opaque zone) for all otolith samples indicates that Pacific bluefin tuna experienced a decreasing water temperature. Age-0 juveniles predominantly inhibit in the surface mixed layer (Kitagawa et al., 2000), and they are known to migrate northward in summer along the coastal regions of Japan and southward for overwintering in the East China Sea and nearshore waters in the Pacific side of Japan (Itoh et al., 2003, Fujioka et al., 2018). The increase of the $\delta^{18}O_{\text{otolith}}$ values thus reflects the actual water temperature change experienced by the individuals from autumn to winter. The increasing patterns of the $\delta^{18}O_{\text{otolith}}$ in the first year of life observed in this study are consistent with the results previously reported by Shiao et al. (2010), with much greater temporal resolutions (days to weeks) with high precision and accuracy achieved by SIMS techniques.

The $\delta^{18}O_{\text{otolith}}$ profiles after the age 0 showed similar seasonal fluctuations, but with much less variation. Figure 3-13 shows monthly averaged water temperature in 2013 at the surface layers (10 to 30 m) in major nursery grounds (Tohoku area and northern Sea of Japan in summer and the East China Sea in winter) of immature bluefin tuna. The estimated water temperature are approximately 8°C higher than the water temperature in respective nursery ground, suggesting that the observed $\delta^{18}O_{\text{otolith}}$ profiles after the age 0 is likely due to the effect of thermoregulatory ability developed by Pacific bluefin tuna. For endothermic fish (warm-bodied) such as Pacific bluefin tuna and other tuna species, the $\delta^{18}O_{\text{otolith}}$ of immature and adult stages do not merely reflect ambient (*in-situ*) water temperature but, rather, the elevated, internal body temperatures. Using counter current heat

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exchangers known as *retia mirabilia* (Dickson & Graham, 2004), bluefin tuna have the capacity to elevate the temperature of their viscera, red (slow-twitch, oxidative) myotomal muscle fibers, eyes, and brain (Linthicum & Carey, 1972, Carey & Lawson, 1973). When juvenile bluefin tuna reach about 20.0 cm in fork length at age-0 (about 2 months after hatching) (Kubo et al., 2008), thermoregulatory ability begins to develop and they are able to maintain their body temperature < 1°C above that of the surrounding water (Furukawa et al., 2017). Thermoregulatory ability of Pacific bluefin tuna increases as fish grow and develop. For example, water temperatures of the peritoneal cavity of adult fish (a 250 kg bluefin tuna) could be 10°C higher than ambient temperatures constructed from acoustic telemetry data (Kitagawa et al., 2006). Since temperatures >35°C are lethal for Pacific bluefin tuna, the observed upper temperature range (30°C) in immature and adult stages is likely a result of physiological thermoregulation to avoid overheating. Understanding how thermoregulatory ability changes with body size and its physiological effects on $\delta^{18}O_{otolith}$ will facilitate better interpretation of the temperature data obtained from this technique.

The estimated temperatures experienced during the larval stages ranged between 26.7°C and 30.7°C among the individuals, with the mean temperature of 28.5 ± 3.0 °C (± 2 SD). The field surveys have collected Pacific bluefin tuna larvae in sea surface temperatures (SSTs) between 23.5°C and 29.5°C in the two main spawning grounds (Yonemori 1989, Abe et al., 2014, Suzuki et al., 2014), with higher concentrations of larvae found around 27°C. The estimated temperatures overlap with the range of temperatures observed for larval occurrence of Pacific bluefin tuna. As newly hatched bluefin larvae tend to stay in warm waters within or near the spawning grounds for their optimal growth and survival, temperatures estimated from the core $\delta^{18}O_{\text{otolith}}$ corresponding to the early larval period may also serve as a useful indicator of spawning temperatures or the location of spawning grounds. One of the specimens (T64R) appeared to experience relatively warm temperatures (>30°C) associated with decreased growth rates and survival of larvae in the laboratory (Kimura et al., 2007). Although there is evidence that the eggs of Pacific bluefin tuna hatched to normal larvae at 31.5°C in

rearing experiments (Miyashita et al., 2000), the resulted higher temperature estimated for this specimen could be a result of the high-resolution sampling of SIMS which captured several warmer days. Despite the need to increase sample size to accurately judge whether the estimated temperatures are realistic, particularly for the specimens with higher temperature, our results suggest that SIMS $\delta^{18}O_{otolith}$ analysis coupled with a microvolume CF-IRMS $\delta^{18}O_{otolith}$ analysis and a species-specific temperature-dependent fractionation equation is an effective method for reconstructing ambient water temperatures experienced by fish and inferring their early life characteristics, which are difficult to obtain with limited resolutions of the conventional methods.

We developed a novel method to estimate ambient temperatures experienced during the larval stage of fish species using SIMS and microvolume CF-IRMS $\delta^{18}O_{otolith}$ analyses. Microvolume $\delta^{18}O_{otolith}$ analysis revealed that the SIMS $\delta^{18}O_{otolith}$ values were 0.41‰ lower on average than CF-IRMS $\delta^{18}O$ values. High-resolution SIMS $\delta^{18}O_{otolith}$ analysis on Pacific bluefin tuna otoliths achieved greater spatial and temporal resolutions with high precision and accuracy compared to the conventional IRMS methods. The $\delta^{18}O_{otolith}$ profiles of all samples showed distinct seasonal variations, reflecting ambient water temperatures experienced by an individual fish. The developed protocol is useful especially for smaller otoliths with narrow growth increments. SIMS $\delta^{18}O_{otolith}$ analysis coupled with micromilling and microvolume $\delta^{18}O_{otolith}$ analysis allows for microscale examinations of otoliths, and more detailed information on the thermal life history of fish can be obtained compared to conventional IRMS methods. This novel method is a powerful tool for the reconstruction of environmental histories of various fish species and has important implications for understanding how ocean warming is potentially affecting the early life history of fish.



Fig. 3-1. Catch locations and landing ports of Pacific bluefin tuna *Thunnus orientalis* collected in this study.



Fig. 3-5. Transverse otolith section of Pacific bluefin tuna *Thunnus orientalis*. Age was determined by counting the number of opaque zones (annual increments) indicated by yellow arrows.



Fig. 3-6. Catch locations of Pacific bluefin tuna *Thunnus orientalis* otolith samples. Two samples (T118R and T131R) were caught in nearshore waters off Kii Peninsula by a small local boat and thus the locations are not shown on the map.



Fig. 3-7. Cross-sectional electron probe micro analyzer (EPMA) image of bluefin tuna otolith with beam spot locations. (b) SIMS beam spot sputtered with a $^{133}Cs^+$ primary ion beam focused to a diameter of 10 to 15 µm. (c) Otolith thin section from Pacific bluefin tuna *Thunnus orientalis* embedded in epoxy resin with a calcite UWC-3 standard. The sample surface was mirror-finished and coated with gold.



Fig. 3-8. Images of Pacific bluefin tuna *Thunnus orientalis* otolith thin section (T64R) used for micromilling. (a) Before milling, (b) after milling unwanted areas to avoid cross-contamination, (c) and after milling the target milling path (the area inside the red line). Black dots indicate the locations of beam spot where SIMS δ^{18} O_{otolith} values are stable (-3.01‰, -2.93‰, -3.04‰, respectively from the top to bottom [VPDB]). The images were taken under an optical microscope for (a), and a stereo microscope for (b) and (c).



Fig. 3-9. Comparison of SIMS and CF-IRMS $\delta^{18}O_{\text{otolith}}$ values (‰, VPDB). (a) CF-IRMS $\delta^{18}O_{\text{otolith}}$ values were significantly higher than SIMS $\delta^{18}O_{\text{otolith}}$ values (*: Wilcoxon signed-rank test, p < 0.001). The bottom and top of the box represent the first (25th percentile) and third (75th percentile) quartiles of the data distribution, respectively. The thick horizontal line inside the box indicates the median, and the upper and lower whiskers represent the maximum and minimum values, respectively.



Fig. 3-10. SIMS $\delta^{18}O_{\text{otolith}}$ values were 0.41‰ lower on average than CF-IRMS $\delta^{18}O_{\text{otolith}}$ values (solid line) when a linear regression with a slope of 1 was fit to the data ($r^2 = 0.78$). Dashed line indicates 1:1. Horizontal and vertical error bars represent CF-IRMS analytical precision (± 1SD) and propagated error of the average SIMS measurements (± 2SD), respectively. Note that one CF-IRMS $\delta^{18}O_{\text{otolith}}$ measurement (grey dot) that had a lower value than the average SIMS $\delta^{18}O_{\text{otolith}}$ value was considered as an outlier based on Tukey's outlier detection method and thus excluded for the calculation of the average SIMS–CF-IRMS difference.



Fig. 3-11. Offset-corrected SIMS $\delta^{18}O_{otolith}$ (‰, VPDB, orange plot on the left y-axis) and reconstructed water temperature (°C, dark blue plot on the right y-axis) of five Pacific bluefin tuna (*Thunnus orientalis*). Error bars represent the spot-to-spot precision of SIMS analyses and the estimated precision of the reconstructed water temperatures (± 2SD). Ambient water temperatures were estimated from the offset-adjusted SIMS $\delta^{18}O_{otolith}$ values using an oxygen fractionation equation for Pacific bluefin larvae proposed by Kitagawa et al. (2013). Dashed vertical lines indicate the locations of annual increments (opaque zone), with numbers showing calendar years of otolith formation. Note $\delta^{18}O_{otolith}$ at the most outer edge was not measured for T104L and T131R, and therefore dashed lines are shown beyond the plots.



Fig. 3-12. Background-corrected ¹⁶OH^{-/16}O⁻ plotted against (a) SIMS δ^{18} O_{otolith} values and (b) SIMS–CF-IRMS δ^{18} O_{otolith} differences measured for 5 Pacific bluefin tuna (*Thunnus orientalis*) otolith samples.

Table 3. Biological information, sampling data, and number of $\delta^{18}O_{\text{otolith}}$ analyses of Pacific bluefin tuna *Thunnus orientalis* used in the present study. Fork length was calculated with a weight-length relationship established by Kai (2007). Two life-history transect lines were analyzed for T64R by SIMS. Numbers in parentheses with asterisks indicate the number of supplementary samples (removed otolith powders that were milled before milling the target milling areas to avoid cross-contamination) analyzed by CF-IRMS.

Sample ID	Catch location	Date of catch	Weight (kg)	Estimated fork length (cm)	Number of SIMS spots	Number of CF-IRMS measurements
T64R	30.0°N, 134.0°E	26 April 2017	146.0	191.1	78 (35 and 43)	5 (2)*
T75R	30.0-32.0°N, 136.0°E	30 April 2017	52.2	136.2	42	-
T104L	27.0-28.0°N, 132.0-134.0°E	18 May 2017	99.2	168.3	53	8 (3)*
T118R	Nearshore off Kii Peninsula	22 May 2017	138.0	187.6	44	9 (4)*
T131R	Nearshore off Kii Peninsula	28 May 2017	118.0	178.2	42	-

Chapter 4 cannot be viewed because the contents will be published in an academic journal.

Chapter 5 cannot be viewed because the contents will be published in an academic journal.

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