3 Results

3.1 Results of U/Th dating in Ishigaki Island

3.1.1 Reliability of dating accuracy for samples

My ²³⁰Th dating results, together with the previously reported ¹⁴C ages, are shown in Tables 1 and 2. The δ^{234} U_{initial} values are used as an indicator of diagenetic alteration of fossil corals. If the initial δ^{234} U values of the samples are similar to those of modern corals and seawater, the samples have been kept in the closed system regarding uranium and thorium. In this study, all of our δ^{234} U_{initial} measurements of *Porites* corals were similar to δ^{234} U value of modern corals (145.8 ‰) reported by Cheng et al. (2000). By using the new ²³⁴U decay constant value, the δ^{234} U value of modern corals can be calculated to be 144.1 ‰, which value is still close to our δ^{234} U values. There are some other studies. Delanghe et al. (2002) reported 146.6 ‰ for the δ^{234} U value of modern coral, while Delanghe et al. (2002) and Robinson et al. (2004) reported that value of modern seawater as 149.6 and 146.6 ‰, respectively. These are all similar to our δ^{234} U_{initial} measurements. Overall, these indicate that correction for such alteration is probably not needed.

The ²³⁰Th ages of two horizons of one modern coral core (Samples 98IY03-L and 98IY03-P) and those of three horizons of one *Porites* boulder core IYT1 (Samples IYT1-6-3, IYT1-11-3 and IYT1-16-3) generally agreed well with the ages determined by annual growth band counting (Table 1; Figure 8). However, in detail, the age differences for two samples 98IY03-L and IYT1-6-3 were larger than the estimated error from analytical precision (less than 4 years). Possible cause of the age disagreements is probably heterogeneous inclusion of detrital materials into coral skeleton. The ²³⁰Th/²³²Th ratio of detrital materials may be different from that of dissolved phase in seawater, but the difference becomes significant only for young modern samples (Shen et al. 2008). Thus, these occasional biases from true age would be insignificant for those coral boulders although replicate measurements of coeval samples may be further required to verify the ²³⁰Th age accuracy.

	ages and 1	Table 1. C ages and 111 ages of inassive <i>F or ues</i> coral ocurates on isingaki island.	NO TOLINES CO			minist min						
	Loc	Location	Size	^{14}C	¹⁴ C age	$\delta^{234} U_{Initial}{}^{a}$	a Initial	²³⁰ Th Age ^b	Λge^{b}	Date obtained by counting bands $^{\circ}$	$\Delta \mathbf{R}$	Median revised corrected ¹⁴ C age
sample ID	latitude	longitude	[m]	[year]	[year BP $\pm \sigma$]	$[\%_0 \pm 2\sigma]$: 2o]	$[A.D. \pm 2\sigma]$	Ε 2σ]	[A.D.]	[years]	[A.D.]
Coral boulders	ulders											
GPS65	24°30.024'	124°17.043'	$4 \times 5 \times 5$	520	± 30	147.6	± 1.3	1770	±3		-16	1764
IYT1-1-3	24°32.820'	124°18.666'	$4 \times 3 \times 3$	550	± 30	145.4	± 1.6	1780	1 ∓		16	1766
IYT4-1-2	24°33.971'	124°19.809'		620	± 30	147.9	± 1.3	1772	±5	-	84	1764
IYT3-1	24°33.650'	124°19.455'	ı	670	±30	147.2	±1.2	1627	±3		-36	1603
IYT2-1	24°33.550'	124°19.357	,	720	± 30	145.4	± 1.2	1607	Έ	-	5	1594
$GPS72^{d}$	24°31.199'	124°17.968'	$7 \times 8 \times 4$	720	± 30	145.8	±1.1	1601	± 2		9	1595
GPS83	24°26.433'	124°15.427	$3 \times 3 \times 2$	500	±150	ı	ı	ı	,		ı	1837
GPS114	24°26.050'	124°15.044'	$3 \times 3 \times 2$	510	±80	ı	ı	ı	,		ı	1857
GPS122	24°23.839'	124°15.208'	$3 \times 4 \times 3$	560	±80		,	ı	,		,	1786
GPS116	24°26.054'	124°15.061'	$3 \times 3 \times 2$	600	±80	,	,	ı	,		,	1756
GPS103	ı	ı	$2 \times 2 \times 2$	650	± 80		,	ı	,	-	,	1621
GPS115	24°26.065'	24°26.065' 124°15.037'	$3 \times 3 \times 1$	720	± 80		,					1581
Modern coral core	yral core											
98IY03-L	'		'	425	± 30	144.7	± 1.3	1904	± 2	1910-1912	-25	ı
98IY03-P		ı		485	±95	146.9	± 1.2	1855	± 2	1852-1856	2	
Fossil coral core	al core											
IYT1-6-3				•		148.3	± 1.2	1705	44	1711-1713	ı	
IYT1-11-3	'			•		147.0	± 1.3	1659	± 2	1658-1659		ı
IYT1-16-3		ı		ı	ı	147.9	± 1.1	1606	± 2	1603-1606	ı	·
Decay consta: ^a δ ²³⁴ U _{initial} w(^b Corrected ²³ ^c Obtained by ^d It is difficult	nts $\lambda_{230} = 9.17$ as calculated t ^o Th ages assu counting ann t to evaluate th	Decay constants $\lambda_{230} = 9.1705 \times 10^{-6}$ year ⁻¹ and $\lambda_{234} = 2.82206 \times 10^{-6}$ year ⁻¹ ; $\lambda_{238} = 1.55125 \times 10^{-10}$ year ⁻¹ (Jaffey et al. 1971). ^a δ^{234} U _{mial} was calculated based on the ²³⁰ Th age (T), i.e., δ^{234} U _{measued} × $e^{\lambda^{234} \times T}$. δ^{234} U = (234 U) ²³⁸ U _{lactivity} - 1) × 1000. ^b Corrected ²³⁰ Th ages assuming an initial ²³⁰ Th ²³² Th atomic ratio of 4.4 ± 2.2 × 10 ⁻⁶ , which is the values for a material at secular ^c Obtained by counting annual growth bands from the top of the coral.	⁻¹ and $\lambda_{234} = 2$ ^b Th age (T), i ²³⁰ Th/ ²³³ Th at ds from the to ion of this bo		< 10^{-6} year- initial = δ^{234} U io of 4.4 ± 2 coral.	1, $\lambda_{238} = 1.5$ measured × $e^{\lambda z}$ 2 × 10 ⁻⁶ , w	5125×10^{-1} $34 \times T$, $\delta^{234}U$ which is the	10^{-6} year ⁻¹ ; $\lambda_{238} = 1.55125 \times 10^{-10}$ year ⁻¹ (Jaffey et al. 1971). 10^{-6} year ⁻¹ ; λ_{238} U _{measured} × e^{2234} V = $(1^{234}$ U/ 238 U _{metriviy} - 1) × 100 of 4.4 ± 2.2 × 10 ⁻⁶ , which is the values for a material at securities or al.	uffey et al. J] _{activity} − . material	Decay constants $\lambda_{230} = 9.1705 \times 10^{-6}$ year ⁻¹ , and $\lambda_{234} = 2.82206 \times 10^{-6}$ year ⁻¹ , $\lambda_{238} = 1.55125 \times 10^{-10}$ year ⁻¹ (Jaffey et al. 1971). ^a δ^{234} U _{mial} was calculated based on the ²³⁰ Th age (T), i.e., δ^{234} U _{messued} × $e^{\lambda^{234} \times T}$. δ^{234} U = (f^{234} U) ²³⁸ U _{mativity} - 1) × 1000. ^b Corrected ²³⁰ Th ages assuming an initial ²³⁰ Th ²³² Th atomic ratio of 4.4 ± 2.2 × 10 ⁻⁶ , which is the values for a material at secular equilibrium with the bulk earth ²³² Th ²³⁸ U value of 3.8. ^c Obtained by counting annual growth bands from the top of the coral.	earth ²³² Th/	²³⁸ U value of 3.8.

Table 1. ¹⁴C ages and ²³⁰Th ages of massive *Porites* coral boulders on Ishigaki Island.

Sample ID	²³⁸ U [ppb]	D [9	²³² Th [ppt]	t] t]	230 Th / 232 Th [atomic × 10 ⁻⁶]	232 Th $\times 10^{-6}$]	δ ²³⁴ U ^a [‰, measured]	U ^a ısured]	230 Th / 238 U [activity × 10 ⁻⁶]	$/ ^{238}$ U · × 10 ⁻⁶]	²³⁰ Th Age [year, uncorrected]	²³⁰ Th Age r, uncorrected]	²³⁰ Th Age ^b [year, corrected]	Age ^b rrected]	$\delta^{234} U_{\text{Initial}}^{\circ}$ [%0, corrected]	rected]
Coral boulders	ders															
GPS65	2549	± 3	190	± 4	557	± 13	147.5	± 1.3	2527	±25	240	± 2	238	±3	147.6	± 1.3
IYT1-1-3	2498	± 3	35	± 8	2834	± 650	145.3	± 1.6	2397	±14	228	±1	228	±1	145.4	± 1.6
IYT4-1-2	2709	± 3	723	± 14	158	± 3	147.8	± 1.3	2555	±24	243	± 2	236	± 5	147.9	± 1.3
IYT3-1	2591	± 3	243	± 5	708	± 15	147.0	± 1.2	4026	± 22	383	± 2	381	± 3	147.2	± 1.2
IYT2-1	2759	± 3	961	± 19	203	±4	145.2	± 1.2	4298	±22	410	± 2	401	十7	145.4	± 1.2
GPS72	2861	± 3	196	4 ±	1033	± 21	145.6	± 1.1	4288	± 20	409	± 2	407	± 2	145.8	± 1.1
Modern coral core	l core															
98IY03-L	2529	± 3	58	± 1	<i>L</i> 6 <i>L</i>	±23	144.7	± 1.3	1101	±22	105	± 2	104	± 2	144.7	± 1.3
98IY03-P	2367	± 2	83	± 2	763	± 19	146.9	±1.2	1621	± 21	154	± 2	153	± 2	146.9	± 1.2
Fossil coral core	core															
IYT1-6-3	2427	± 3	84	± 2	1511	±39	148.2	± 1.2	3190	± 42	303	± 4	303	± 4	148.3	± 1.2
IYT1-11-3	2356 ±2	± 2	56	± 1	2554	± 56	146.9	± 1.3	3669	± 23	349	± 2	349	± 2	147.0	± 1.3
IYT1-16-3	2187	± 2	46	± 1	3293	± 71	147.7	± 1.1	4225	± 21	402	± 2	402	± 2	147.9	± 1.1

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^a $\delta^{234}U = ([^{234}U/^{238}U]_{activity} - 1) \times 1000.$

^b Corrected ²³⁰Th ages assuming an initial ²³⁰Th/²³²Th atomic ratio of $4.4 \pm 2.2 \times 10^{-6}$, which is the values for a material at secular equilibrium with the bulk earth ²³²Th/²³⁸U value of 3.8. ^c δ^{234} U_{minial} was calculated based on the ²³⁰Th age (T), i.e., δ^{234} U_{menued} $\times e^{\lambda234 \times T}$.

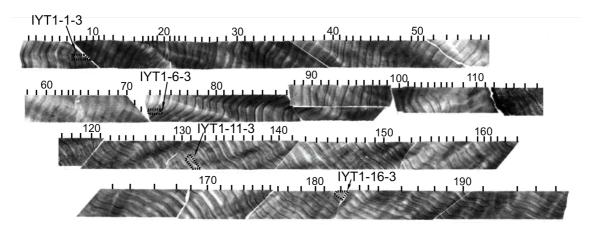


Figure 8. Positive X-radiograph prints of 6-mm-thick slabs of the *Porites* boulder core IYT1. Alternating bands of high (dark color) and low (light color) skeletal density are visible. A high/low density band pair represents one year of growth. The number of years of coral growth correspond to that of density band pairs. Skeletal portions used for ²³⁰Th dating are outlined by dotted lines. Core length is 3.2 m and core diameter is 5.5 cm.

3.1.2 Correlation between dating results and historical documents

The "Kyuyo" and "Yaeyama-jima Nenraiki" chronicles describe three inundation disaster events on Ishigaki Island: the 1771 Meiwa earthquake tsunami and two other inundation events with unspecified causes (1625, 1714). These inundation events could be caused not only by tsunamis but also by extreme storms. For these two inundation events, the "Yaeyama-jima Nenraiki" chronicles report only that "a large wave came from", so that it is difficult to know whether to ascribe them to an earthquake tsunami or a severe typhoon. In addition, a tsunami that occurred on 20–21 October in 1687 as a result of earthquakes off the Peruvian coast (two large earthquakes with magnitudes estimated at 8.0 and 8.4, Dorbath et al. 1990) and was recorded in Okinawa Island, though not in Ishigaki Island. Although no damage was described in the chronicles, this tsunami should have struck Ishigaki Island. Therefore, at least four recognized coastal paleo-inundation events affected Ishigaki Island during the 17–19th centuries. Of the six *Porites* boulders dated by the ²³⁰Th method (Figure 2b-2f), two (samples GPS65 and IYT4-1-2) could be directly related to the 1771 Meiwa earthquake tsunami within the analytical error (2σ) of approximately 5 years (Table 1). The date of another specimen (sample IYT3-1) corresponds to the 1625 event (Table 1).

3.1.3 Calculation of local marine reservoir correction values

I plotted the marine reservoir ages and ΔR values calculated from pairs of ²³⁰Th and ¹⁴C ages obtained from the same coral samples as a function of age (Figure 9). I calculated mean ΔR values

for three time windows: 10 ± 37 years for 1600-1780; -8 ± 20 years for 1600-1630; and 28 ± 42 years for 1770-1780. The revised conversions of the ¹⁴C ages to calendar ages based on the newly determined ΔR values are shown in Table 1.

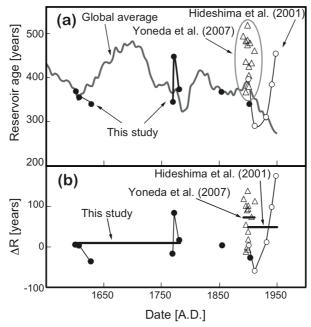


Figure 9. Temporal variations of (a) radiocarbon reservoir ages and (b) ΔR in the Kuroshio region. Symbols: closed circles, this study; open circles, Hideshima et al. (2001); open triangles, Yoneda et al. (2007). The gray curve in panel (a) shows the mean global average of marine reservoir ages (MARINE04; Hughen et al. 2004). The horizontal bars in panel (b) indicate the mean ΔR values of each study: this study, 10 ± 37 years (n = 6); Hideshima et al. (2001), 49 ± 41 years (n = 5); Yoneda et al. (2007), 73 ± 17 years (n = 14).

3.2 Results of Radiocarbon darting in Southern Ryukyu Islands

3.2.1 Reliability of dating accuracy for samples

To check the accuracy of radiocarbon dating, we measured ¹⁴C ages of different surfaces of one colony (sample MH11; Figure 6) and the results were well agreed within the dating error (Table 3). Therefore, if the youngest parts of the boulders were collected exactly, accurate sample ages would be obtained by ¹⁴C dating. Reproducibility of calibrated ¹⁴C ages was also evaluated by measuring same sample (sample MH11-1; Figure 7) for several times (Table 4).

I measured ¹⁴C ages of boulders found at Miyako (sample MH5 and MH10; Figure 10) and Ishigaki Islands (sample H1), which have different growth axes in the process of growing (Table 5). If the growth rate of the coral are assumed as 1 centimeter per a year in Ryukyu region, difference of the ages were consistent, indicating that these boulders have a potential for recording the tsunami events twice (Table 5).

	Loc	ation		Size		¹⁴ C	age	Calit	prated ¹⁴ C age [cal AI	D]
Sample ID	Latitude	Longitude		[m]		[year H	$BP\pm\sigma]$	±1σ range (probability)	±2σ range (probability)	Median
MH11-1-1						425	±57	1841 — 1951	1722 — 1951	1877
MH11-2	24°43.744'	125°27.002'	$2.5 \times$	2.2 ×	1.5	460	± 48	1827 — 1951	1721 — 1951	1904
MH11-3						438	± 71	1819 — 1951	1714 — 1951	1755

Table 3. The ¹⁴C ages of the surface parts of massive *Porites* boulder (MH11).

Table 4. Reproducibility of calibrated ¹⁴C ages by dating same sample (MH11-1).

	Loc	ation		Size		¹⁴ C	age	Calib	orated ¹⁴ C age [cal AI	ומ
Sample ID	Latitude	Longitude		[m]		[year H	$BP \pm \sigma$]	±1σ range (probability)	±2σ range (probability)	Median
MH11-1-1						425	±57	1841 — 1951	1722 — 1951	1877
MH11-1-2	24°43.744'	125°27.002'	2.5 ×	2.2 ×	1.5	380	±39	1881 — 1951	1809 — 1951	1865
MH11-1-3						583	±53	1667 — 1827	1626 — 1951	1861

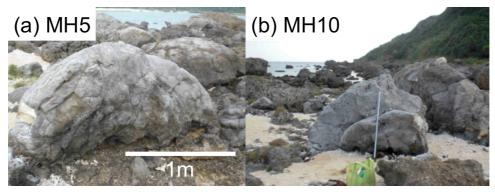


Figure 10. Photographs of Porites boulders having different growth axes at Miyako Island.

	Loc	ation	Size		¹⁴ C	age	Calit	orated ¹⁴ C age [cal AI	0]
Sample ID	Latitude	Longitude	[m]		[year I	$BP\pm\sigma]$	±1σ range (probability)	±2σ range (probability)	Median
Boulder MH5									
MH5-1	24°43.744'	125°26.995'	18 × 15	× 1	1184	±52	1173 — 1292	1071 — 1324	1226
MH5-2	27 73./44	125 20.995	1.0 ^ 1.5	~ 1	1125	±104	1182 — 1386	1057 — 1440	1268
MH10-1					1382	±49	965 — 1114	895 — 1183	1034
MH10-2	24°43.724'	125°26.993'	- × -	× -	1327	±106	988 — 1219	849 — 1296	1089
Boulder H1									
H1-1	-	-	- × -	× -	1456	± 62	880 — 1036	781 — 1120	952
H1-2	-	-	- ~ -		1351	±76	985 — 1170	888 — 1254	1069

Table 5. The ¹⁴C ages of the surfaces of massive *Porites* boulders having different growth axes.

As ΔR values are generally not constant over time, strictly it may not be appropriate to apply one estimated ΔR value for 1600-1780 to samples of different ages. However, although no data exist in Ryukyu region except for Hideshima et al. (2001), Yoneda et al. (2000, 2007) and this study. Yamaguchi et al. (2004) and Yu et al. (2010) reported local ΔR values around Taiwan and South China Sea near Southern Ryukyus, respectively, and suggested that the variation of ΔR values were not substantial since middle Holocene. Therefore, in this study I applied my estimated ΔR value to convert ¹⁴C ages of all samples into calendar ages.

3.2.2 Calibrated ¹⁴C ages of Porites coral boulders

I measured the age of 125 samples from 77 boulders (Table 6). The locations of ¹⁴C-dated boulders were shown together with own dates in Figure 11. The results of ¹⁴C measurements also showed various ages from more than 2500 years ago to nearly present (Figure 12), which could be attributed to various paleo-tsunamis other than the 1771 Meiwa tsunami. In Miyako and Kurima Islands, many boulders were attributed to tsunamis older than the Meiwa tsunami (Figure 11). Some *Porites* boulders showed much older ages, which were about 3900, 4600, and 5800 years ago. According to these results, past tsunami disasters were likely happened in this region constantly during this period.

Median 1815 1781 1699 1639 1593 1904 1783 1814 1775 1736 1670 1597 1912 1900
1781 1699 1639 1593 1904 1783 1814 1775 1736 1670 1597 1912
1781 1699 1639 1593 1904 1783 1814 1775 1736 1670 1597 1912
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Table 6. The ¹⁴C ages of 77 *Porites* coral boulders at Southern Ryukyu Islands.

	Loc	ation		Size		¹⁴ C	age	Calib	orated ¹⁴ C age [cal AI	0]
Sample ID	Latitude	Longitude		[m]		[year H	$BP \pm \sigma]$	±1σ range (probability)	±2σ range (probability)	Median
Boulders on Mi	nna Island									
GPS229	24° 45.448'	124° 41.283'	4 >	< 3.6 →	< 2.6	388	± 55	1861 — 1951	1725 — 1951	1893
GPS232	24° 45.464'	124° 41.418'	_	_	_	2502	± 58	-322 — -130	-368 — -29	-205
GPS236	24° 45.264'	124° 41.626'	_	_	_	521	± 50	1714 — 1884	1694 — 1951	1813
GPS237	24° 45.212'	124° 41.697'	1.1 >	< 1.1 →	0.3	3918	± 55	-2037 — -1845	-2126 — -1745	-1939
GPS240	24° 44.836'	124° 41.897'	4.7 >	< 3.9 →	3.4	527	± 50	1710 — 1877	1691 — 1951	1807
GPS241	24° 44.819'	124° 41.899'	2.8 >	< 2.5 →	< 2	494	±47	1774 — 1951	1709 — 1951	1839
GPS242	24° 44.818'	124° 41.901'	3.5 >	< 2.7 →	< 2	693	±52	1545 — 1675	1477 — 1803	1614
GPS243	24° 44.816'	124° 41.899'	2.1 >	< 1.9 →	< 2	684	±46	1556 — 1681	1490 — 1803	1623
GPS249	24° 45.109'	124° 41.463'	_	_	_	1298	±49	1050 - 1182	1001 — 1256	1121
GPS252	24° 45.039'	124° 41.598'	_	_	_	618	±49	1645 — 1808	1540 — 1881	1706
GPS254	24° 44.918'	124° 41.724'	_	_	_	551	±49	1690 — 1851	1672 — 1951	1785
GPS255	24° 44.932'	124° 41.769'	_	_	_	985	± 48	1332 — 1428	1294 — 1465	1381
Boulders on Mi	yako Island									
MH1	24°43.700'	125°26.981'	1.5 >	< 1	_	2568	±42	-351 — -212	-396 — -126	-276
MH2	24°43.710'	125°26.988'	1.8 >	(1.2	_	382	±122	1773 — 1951	1679 — 1951	1847
MH4	24°43.720'	125°26.990'	2 >	< 2 >	< 1	1514	±146	730 — 1034	611 — 1214	893
MH5	24°43.720'	125°26.995'	1.8 >	< 1.5 →	< 1	1125	±104	1182 — 1386	1057 — 1440	1268
MH6	24°43.741'	125°26.990'	4 >	< 2.5 →	< 1.5	683	± 56	1546 — 1686	1489 — 1805	1625
MH10	24°43.724'	125°26.993'	—	—	—	1382	±49	965 — 1114	895 — 1183	1034
MH11	24°43.744'	125°27.002'	2.5 >	< 2.2 →	< 1.5	380	± 39	1881 — 1951	1809 — 1951	1904
Boulders on Shi	mochi Island									
GPS267	24° 50.345'	125° 8.780'	_	_	_	2845	± 55	-743 — -548	-783 — -426	-631
GPS272	24° 50.367'	125° 8.663'	—	-	_	651	± 56	1545 — 1720	1502 — 1840	1663
GPS277	24° 50.419'	125° 8.314'	5.5 >	< 5 >	< 3	459	±23	1851 — 1951	1724 — 1951	1882
GPS278	24° 50.421'	125° 8.340'	5.5 >	< 4.3 >	< 3.6	414	±21	1882 — 1951	1819 — 1951	1904
Boulders on Ku	rima Island									
GPS284	24° 43.393'	125° 14.363'	1.5 >	0.7	_	2211	± 61	70 — 245	-26 — 339	153
GPS286	24° 43.472'	125° 14.379'	1.3 >	0.7	_	1621	± 54	700 — 845	660 — 936	782
GPS288	24° 43.528'	125° 14.390'	1.8 >	< 1.4	_	1894	± 53	445 — 594	377 — 661	520

Table 6. The ¹⁴C ages of 77 *Porites* coral boulders at Southern Ryukyu Islands.

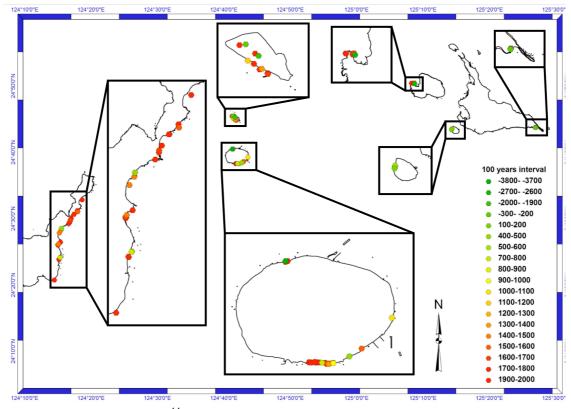


Figure 11. Distribution of the ¹⁴C-dated *Porites* boulders. Color variation shows separating each date with 100 year.

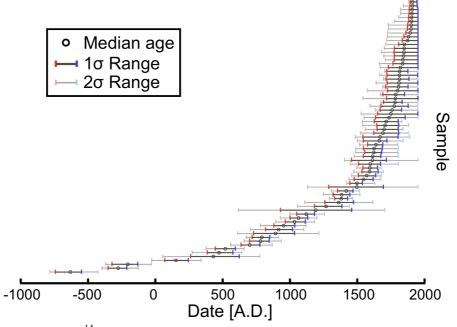


Figure 12. Calibrated ¹⁴C ages of the *Porites* boulders and its age distributions. Some *Porites* boulders showed much older ages, which were not shown in this Figure.